

1        **Enhancing River Channel Dimension Estimation: A Machine Learning**  
2        **Approach Leveraging the National Water Model, Hydrographic Networks,**  
3        **and Landscape Characteristics**

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12       **Key Points:**

- 13       • We present a data-driven approach for predicting channel geometry using a machine  
14       learning (ML) method.
- 15       • ML predictors include National Water Model flow statistics and a suite of landscape  
16       characteristics.
- 17       • Our method exceeds the goodness of fit of its predecessors evaluated across regions of  
18       the United States.

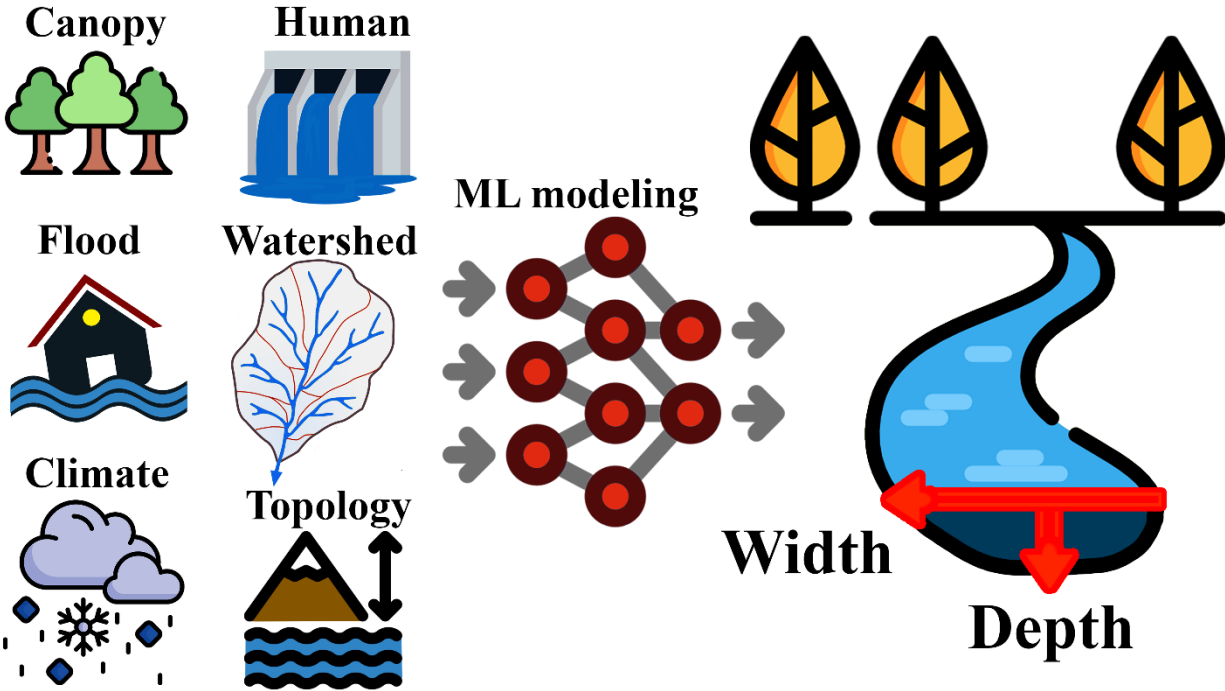
19       **Abstract**

20       Knowledge of bankfull hydraulic geometry represents an essential requirement for  
21       various applications, including accurate flood prediction, hydrological routing, river behavior  
22       analysis, river management and engineering practices, water resource management, and beyond.

23 Our work builds upon an extensive body of literature about estimating bankfull top-width and  
24 depth at ungauged locations to enhance the understanding of observable factors that affect these  
25 parameters. Using more than 200,000 USGS Acoustic Doppler Current Profiler (ADCP) records,  
26 we developed a method employing machine learning (ML) using discharge estimates and  
27 landscape characteristics from sources, including the National Water Model (NWM), the  
28 National Hydrologic Geospatial Fabric network (NHGF), the EPA stream characteristic dataset  
29 (StreamCat), and an array of satellite and reanalysis data products. Our method achieved  $R^2=0.79$   
30 predicting bankfull depth ( $R^2=0.84$  for in-channel conditions) and  $R^2=0.81$  predicting bankfull  
31 top-width ( $R^2=0.8$  for in-channel conditions) in the testing dataset. The depth predictions showed  
32 high skill in plateau regions and low skill in mountainous regions. Our analysis demonstrates the  
33 benefit of data-driven modeling in contrast to other global scaling-based or regional statistical  
34 methods. In summary, our study illustrates how top-width and depth can be better predicted  
35 using ML, reanalysis streamflow simulations, hydrographic networks, and summarized  
36 geospatial data.

37 Keywords: at a station hydraulic geometry, river characteristics, river dimension, hydrofabric,  
38 machine learning, National Water Model

39 **Graphical abstract**



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41

42 **Plain language summary**

43 Accurately estimating (or generalizing) key characteristics of river channels, such as their top-  
 44 width and depth, is valuable for tasks like predicting water flow, modeling water-related  
 45 processes, and mapping flooded areas. Our research builds on existing studies that focus on  
 46 estimating these important channel characteristics and aims to further develop knowledge and  
 47 skills in predicting these channel characteristics. In this work, we use over 200,000 historical  
 48 measurements of channel top-width and depth to develop a machine learning (ML) model to  
 49 estimate channel top-width and depth. The model uses widely available information from the  
 50 National Water Model (NWM) discharge and other datasets that represent land surface  
 51 characteristics, climate, hydrographic connectivity, and human-related structures. The developed  
 52 model performs well compared to other global, regional, and ML-based methods in the literature  
 53 within the Continental United States. Validation of the models across different regions indicated

54 better performance in flatter regions and lower performance in steeper areas. In conclusion, the  
55 study highlights the advantages of using ML techniques to estimate channel geometry more  
56 accurately, paving the way for improved predictions in unmeasured channels.

## 57 **Introduction**

58 The use of accurate estimates of channel bankfull depth and top-width improves channel  
59 flow routing models (Bindas et al., 2024; Brackins et al., 2021a; Getirana et al., 2013; Han et al.,  
60 2020). These dimensions inform generalized cross-sections in large-scale models, and their  
61 adequacy can influence hydrological forecasting (Brackins et al., 2021b; Brakenridge et al., 2012;  
62 Cohen et al., 2019; Heldmyer et al., 2022) and products that rely on them, like flood map  
63 generation (Alfieri et al., 2018; Cohen et al., 2018; Johnson et al., 2019).

64 Apart from large-scale modeling, precise estimates of bankfull depth and top-width  
65 improve flood risk analysis and mapping. These accurate estimations act as a proxy, enhancing the  
66 representation of channel volume below standard DEM elevations. (Bates & De Roo, 2000;  
67 Sichangi et al., 2018; Yamazaki et al., 2009).

68 The interest in estimating bankfull depths and top-widths has a long history in academic  
69 literature. To date, there have been efforts to present global equations that establish a relation  
70 between discharge and bankfull top-width or depth following the geomorphic relation proposed  
71 by Leopold and Maddock (1953). One of the earliest of these efforts was proposed by Moody and  
72 Troutman (2002) for global channels (Eqs. 1 and 2).

$$w = 7.2Q^{0.5 \pm 0.02} \quad (1)$$

73

$$d = 0.27Q^{0.3 \pm 0.01} \quad (2)$$

74 where  $w$  and  $d$  are bankfull top-width and depth, respectively, and  $Q$  is the discharge.

75 Frasson et al. (2019) proposed an alternative drainage area relation to estimate channel  
76 top-width, hypothesizing and affirming that channel top-width is directly associated with  
77 catchment area and channel meander wavelength. They aimed to regionalize the above-  
78 mentioned global relations by using the  $Q/A$  ratio, where  $A$  represents the total drainage area and  
79  $Q$  is discharge obtained from the Global Runoff Data Center (GRDC). This approach resulted in  
80 a global dataset (based on HydroBASIN; Lehner & Grill, 2013) of estimates for bankfull top-  
81 width and depth. They subsequently conducted a comparison of these estimates with a set of  
82 bankfull top-width records derived from Landsat data. However, these equations are limited to  
83 channel reaches below  $60^\circ\text{N}$  and top-width greater than the 30m resolution of Landsat scenes. In  
84 their validation, they found errors ranging from 8 to 62%. Other researchers have used these  
85 equations to compute bankfull estimates of channel top-width and depth to support routing  
86 attributes in land surface models (Han et al., 2020; Schumann et al., 2013).

87 Bieger et al. (2015) proposed regional equations based on regression for different  
88 physiographic divisions across the United States and found that annual precipitation and  
89 temperature provide additional information that improves channel top-width and depth  
90 predictions considerably. A subsequent study by Blackburn-Lynch et al. (2017) developed  
91 regional relations for all Hydrologic Landscape Regions (HLR) and physiographic provinces and  
92 reported higher goodness of fit (GOF) of the discharge-based relations in comparison to drainage  
93 area-based relations. These estimates were used in the development and deployment of WRF-  
94 Hydro and its implementation as the NOAA National Water Model. Recently, Neal et al. (2021)

95 addressed the challenge of missing channel bathymetry more explicitly by demonstrating that  
96 improved bed estimations derived from the simplified, gradually varied flow method  
97 significantly influenced the dynamics of floodplain inundation and storage during minor flood  
98 events.

99 ML models provide a superior alternative to simple regression models by efficiently  
100 learning from multidimensional and complex data, capturing non-linear relationships, and  
101 adapting to diverse feature types (Shen, 2018). ML models tailored for learning specific  
102 hydraulic or hydrologic variables, like channel dimensions, can be trained effectively using  
103 hydrographic datasets such as the National Hydrography Dataset (McKay et al., 2012), the  
104 National Hydrographic Geospatial Fabric (Blodgett et al., 2023; Bock et al., 2022) and its  
105 derived products (Johnson, 2022), Multi-Error-Removed Improved-Terrain (MERIT) Hydro (Dai  
106 Yamazaki et al., 2019), or the Surface Water and Ocean Topography (SWOT) Mission River  
107 Database (SWORD; Altenau et al., 2021), to name a few. Estimating channel characteristics on  
108 networks such as these provides a means to support a range of hydroscience use cases. In 2023,  
109 machine learning applications were introduced by Doyle et al. (2023), which further refined  
110 channel top-width and depth estimates within the United States. They achieved these refined  
111 estimates by using a random forest model parameterized with the Watershed (Ws) summaries  
112 sourced from the EPA StreamCat attributes (Hill et al., 2016). Their method demonstrated the  
113 value of using watershed-based predictors to estimate channel dimensions. However, when using  
114 a large suite of 96 predictors, high dimensionality (referred to as the 'curse of dimensionality';  
115 Köppen, 2000) and high correlation between variables, such as population and housing density  
116 (~ 100%), create collinear predictors within the model, impairing its interpretability post-training  
117 (Chan et al., 2022), and make it less extensible. These challenges can lead to model confusion

118 and a distorted representation of the actual responses, even in ML algorithms such as random  
119 forest methods (Ghahremanloo et al., 2021).

120         Aside from statistical and ML predictions, remote sensing is an alternative solution to  
121 estimate channel dimensions. There have been numerous studies of automated channel top-width  
122 extraction using satellite imagery (Durand et al., 2009; Golly & Turowski, 2017; Monegaglia et al.,  
123 2018; Pavelsky & Smith, 2008; Schwenk et al., 2017). In 2017, Isikdogan et al. (2017) developed  
124 the RivaMap software that automates the extraction of continental-scale river centerline and top-  
125 width for North American rivers using Landsat imagery, and more recently, the RivWidthCloud  
126 software using cloud computing (Google Earth Engine) to extract channel top-widths from a vast  
127 archive of Landsat imagery (Yang et al., 2019). Global Surface Water Explorer (Pekel et al., 2016)  
128 does not provide direct information about channel top-width, but its historical probability map of  
129 water occurrence can be used to distinguish riparian floodplain top-width, bankfull top-width, and  
130 in-channel top-width.

131         Concerning channel depth, satellites like ENVISAT and JASON can provide information  
132 on water surface elevation through altimeter measurements (Kouraev et al., 2004). The recently  
133 launched Surface Water and Ocean Topography (SWOT) satellite shows potential for improving  
134 space-based estimates of channel discharge globally (Durand et al., 2020; Emery et al., 2016).  
135 SWOT capability to measure water surface elevation can be used to capture variations in channel  
136 depth that cannot be directly measured from other remote sensing products such as MODIS,  
137 Landsat, and Sentinel products (Pavelsky et al., 2014). The major constraints on all remote sensing  
138 approaches are (1) the spatial resolution of the data, (2) the quality of the data (e.g., scan lines and  
139 cloud cover), and (3) the computational costs associated with image processing at regional to  
140 continental domains. In the case of (1) the 30-meter resolution of Landsat products and algorithms

141 limit estimation to channels greater than 50 m in top-width at the time of imaging (Andreadis et  
142 al., 2013). Likewise, changes in water surface elevation are constrained by the 50 m resolution of  
143 SWOT, limiting the observable channels to those with top-widths exceeding 100 meters. (Baratelli  
144 et al., 2018). While these estimates are critical for the major river systems, supplementary  
145 information about the tributary systems that feed them is needed for a wide range of engineering,  
146 modeling, and design purposes.

147         In this research, we test the hypothesis that a meta-learning (ensemble ML) methodology  
148 informed by National Water Model (NWM) simulated flow characteristics (such as 100-0.1%  
149 annual exceedance probability discharges) and a suite of land surface and climate variables can  
150 predict bankfull top-width and depth with GOF equal to or better than previous methods. We  
151 hypothesize that the incorporation of NWM simulated flow characteristics and a suite of land  
152 surface and climate variables can significantly improve bankfull top-width and depth predictive  
153 capability. The rest of the paper is as follows: the methods section will describe the input data  
154 used and the model training procedure. The results will discuss the model outputs in relation to  
155 the observed data as well as existing global and regional curves and ML approaches. The  
156 discussion will highlight areas for improvement, applications of this dataset, and the advantages  
157 of taking a hydrofabric-centric approach grounded in evolving federal and international efforts.



## 158 2. Methods

### 159 2.1. Data

#### 160 2.1.1 Observations

161 In this study, we use the HYDRoacoustic dataset in support of Surface Water Oceanographic  
162 Topography (HYDRoSWOT; Bjerklie et al., 2020; Canova et al., 2016). This dataset is  
163 composed of 200,000+ Acoustic Doppler Current Profiler (ADCP) measurements collected at  
164 10,081 USGS stream gauging locations in the United States. From this, we use the recorded  
165 depth, top-width, velocity, and discharge for each campaign.

#### 166 2.1.2. Data Filtering

167 With multiple, time-varying observations for the same location, statistical relationships  
168 can describe the relationship between streamflow, depth, top-width, and velocity. These  
169 relationships have traditionally been described using At a station Hydraulic Geometries (AHG;  
170 Leopold & Maddock, 1953; Shen et al., 2016).

171 The AHG relations are described as:

$$172 \quad Q = TW \times Y \times V = aQ^b \times cQ^f \times kQ^m = ack \times Q^{b+f+m} \quad (3)$$

173 Where TW is the top-width, Y is the depth, and V is the velocity at a recorded streamflow Q  
174 value. Therefore, individual relationships can be described as follows:

$$175 \quad TW = aQ^b \quad (4)$$

$$176 \quad Y = cQ^f \quad (5)$$

177  $V = kQ^m$  (6)

178 and by definition:

179  $a \times c \times k = 1$  (7)

180  $b + f + m = 1$  (8)

181 Observations of hydraulic data, particularly over long periods, are inherently noisy. To reduce  
182 this noise, we leveraged the AHGestimation R package (Johnson et al., 2024) [Click or tap here to](#)  
183 [enter text](#).which provides the ability to filter data based on statistical outliers prior to using an  
184 ensemble-based fitting method to ensure mass is conserved (Eqs 7-8) and error is minimized.

185 Profiles were only kept if (1) they were made between 2000-2015 (2) they had a depth less  
186 than 65 m (corresponding to the Hudson River - the deepest river in the United States) (3) they  
187 had a channel top-width of less than 4 km (corresponding to the widest parts of the Mississippi  
188 River) (4) their recorded discharge values are in agreement with min and max NWIS records of  
189 that site (5) no negative discharge, width, depth, and velocity records. Once reduced, any site that  
190 demonstrated an inverse relationship between discharge and depth or had less than 5 profiles  
191 were removed. After filtering each site, AHGestimation was used to fit equations 4, 5, and 6 to  
192 the filtered HYDRoSWOT data. The AHGestimation (Johnson et al., 2024) uses a combination  
193 of ordinary least square, a nonlinear least square, and a genetic algorithm to fit data while  
194 ensuring mass preservation (equations 7 and 8). From these fits we calculate the coefficient of  
195 determination ( $R^2$ ). We took an arbitrary threshold of 0.6 that can explain more than half the  
196 variability in HYDRoSWOT data and retained sites with a  $R^2 > 0.6$  from AHG fit. In total,  
197 4,229 of the 6,226 initial sites were retained.

### 198 2.1.3. Bankfull and In-Channel top-width and Depth

199 To define in-channel and bankfull discharge at the selected HYDRoSWOT sites, we used  
200 the widely accepted definitions of discharge at 100 and 50% annual exceedance probability,  
201 respectively (Andreadis et al., 2013; Rosgen, 1994; Wilkerson, 2008; Woodyer, 1968). While  
202 recognizing that bankfull and in-channel flow vary across river reaches and correlate with  
203 different flood recurrence intervals, we simplify our approach by labeling 100 and 50% annual  
204 exceedance probability as in-channel and bankfull, respectively. We hypothesized that the 100%  
205 annual exceedance probability discharge corresponds to the absence of bathymetry data, wherein  
206 a digital elevation model (DEM) generates a flat bottom unable to penetrate water. Subsequently,  
207 the 50% annual exceedance probability discharge represents the next higher flood condition, a  
208 widely recognized term in the literature concerning the modeling of bankfull width (Andreadis et  
209 al., 2013).

210 Using historical daily NWIS discharge records (retrieved from DeCicco et al., 2023) at  
211 each HYDRoSWOT site, we computed an annual maxima series and assumed the series may  
212 follow either a generalized extreme value (GEV), generalized Pareto (GP), Log-Pearson Type III  
213 (LP3), or generalized gamma distribution (Metzger et al., 2020; Zhang et al., 2021). To identify  
214 the distribution that best describes the underlying data, the Kolmogorov-Smirnov (KS) method  
215 was used (Ahmad et al., 1988). The selection of the most suitable distribution was determined  
216 through the KS test, which involved computing the KS statistic between the candidate  
217 distribution and the empirical distribution derived from the observed data. Results showed the  
218 majority of sites followed a GEV and GP distribution. Subsequently, the distribution yielding the  
219 smallest KS statistic was identified as the best-fitting model and then used to compute bankfull  
220 and in-channel discharge. The top-width and depth were obtained using the defined AHG

221 relation. Figure S1 shows several HYDRoSWOT sites with corresponding bankfull and in-  
222 channel computed discharges.

223 The river network used in this study is the National Hydrologic Geospatial Fabric (NHGF;  
224 Bock et al., 2022), which is derived from the NHDPlusV2 with modifications to the topology  
225 and network characteristics based on feedback from a collection of federal agencies (Blodgett et  
226 al., 2023). Using this network, we compiled a list of possible predictors that could be used to  
227 explain variability in channel geometry. A list of all considered predictor variables is included in  
228 Table S1. These variables were selected based on literature and are composed of hydraulic,  
229 hydrologic, and climatological characteristics that affect channel dimensions.

230 Following the work of Doyle et al. (2023) and Blackburn-Lynch et al. (2017), we look to  
231 define a suite of catchment and watershed-level characteristics. Like those efforts, we collected a  
232 range of landscape characteristics for all river reaches across the United States using the EPA  
233 StreamCat data set (Hill et al., 2016). This dataset includes information on dams, land use,  
234 climate, hydrology, geology, and more. The outlet features of each total drainage area were  
235 aligned with the reference fabric identifier. In total, we used 58 watershed-level predictors from  
236 this dataset. In addition to these precomputed variables, a set of soil, landscape, and weather data  
237 was obtained and aggregated to the catchment level using climateR (Johnson & Clarke, 2019).  
238 This includes data from TerraClimate (Abatzoglou et al., 2018), POLARIS (Chaney et al., 2019),  
239 NLDAS, GLDAS (Rodell et al., 2004), USGS 3DEP, Leaf Area Index, and the Moderate  
240 Resolution Imaging Spectroradiometer (MODIS) mission. In total, we built 31 predictors from  
241 these sources.

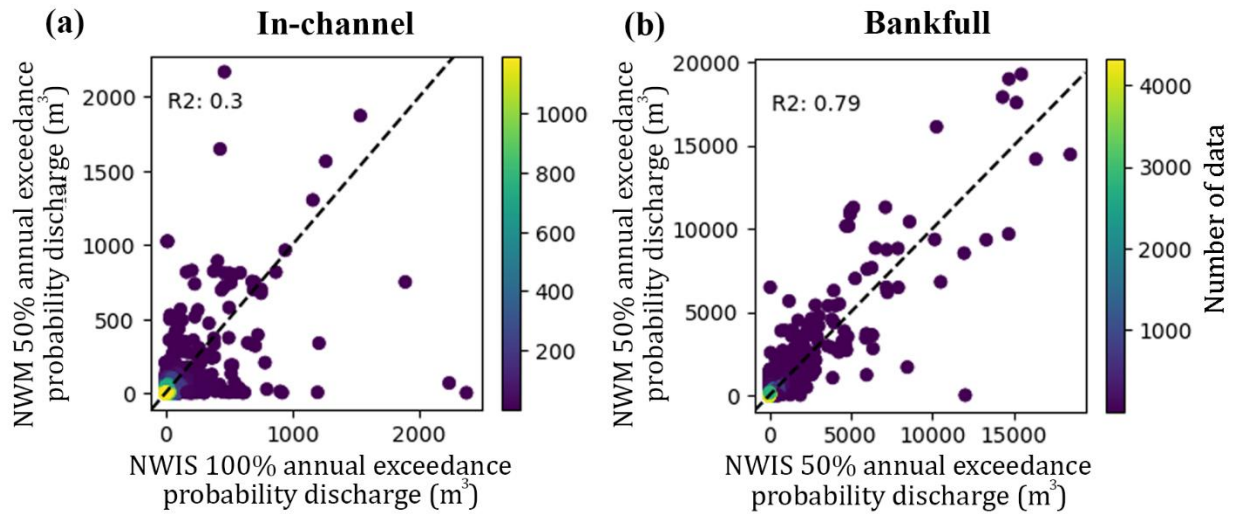
242 Unlike prior efforts, we wanted to explore the impact of adding network connectivity and  
243 streamflow statistics to our predictors. Hydrographic information was taken from the NHGF, and

244 in total, 10 predictors were used that represent general catchment and streamflow characteristics  
245 (see Table S1; source: Reference Fabric).

246 Streamflow statistics were generated for all NHGF reaches from the NWM v2.1 retrospective  
247 simulation (Johnson et al., 2023) and include 20-0.1% annual exceedance probability discharge  
248 (log Pearson Type 3) and the minimum, 25th, 50th, 75th, and maximum flow percentiles. As this  
249 study employed modeled flows, it is essential to note that these are regarded as "synthetic flow  
250 percentiles" due to discrepancies with the gauge measurements. In the training phase, although  
251 all stations contained floods from both NWIS and NWM, we randomly split sites such that half  
252 only contained NWIS bankfull, or in-channel flood, and the other half only contained NWM  
253 bankfull, or in-channel flood, and introduced a binary variable as an indicator of the absence of  
254 NWIS records. This allows the model to predict top-width and depth using accurate NWIS  
255 observations when available and NWM model data otherwise. During the testing phase, bankfull  
256 or in-channel floods are only derived from NWM.

257 As a first-pass quality check, Figure 1 plots the NWM and NWIS one- and two-year flow  
258 estimates against each other to assess the skill of the NWM in these lower flood recurrence  
259 intervals. While it has been shown that the NWM has less skill in predicting low flows as  
260 opposed to high flows (Fang et al., 2023; Johnson et al., 2023), here we test if the addition of  
261 greater floods as predictors compensates for this. Notably, the two-year flows show nice  
262 agreement ( $R^2 = 0.79$ ) with NWIS, while, as expected, the one-year flows show less ( $R^2 = 0.3$ ).

263

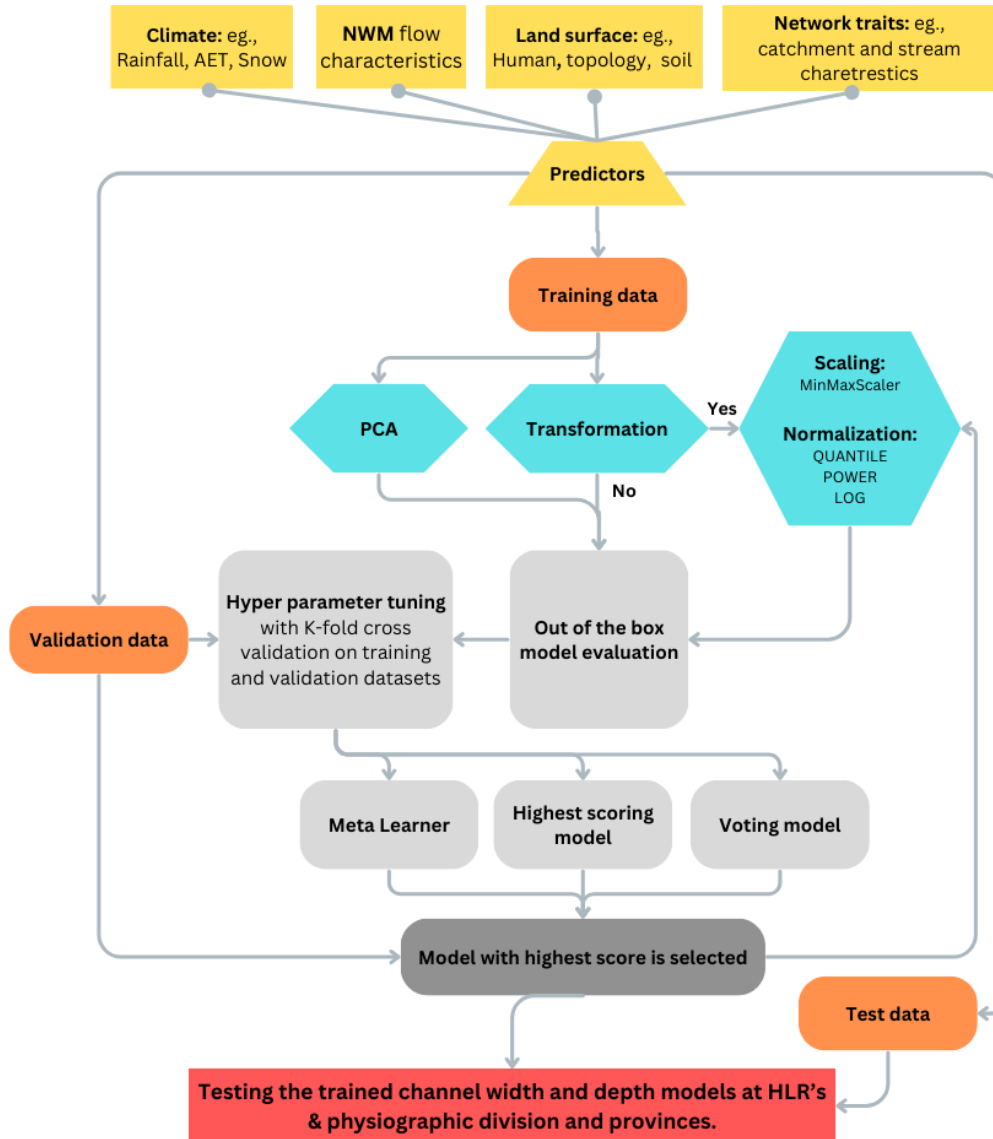


264

265 **Figure 1.** Comparison of NWM and NWIS flood discharges. (a) In-channel or 100% annual  
 266 exceedance probability flow comparison; and (b) bankfull or 50% annual exceedance probability  
 267 comparison.

268 2.2. Modeling

269 We trained four different machine learning models to predict top-width and depth for both  
 270 bankfull (50% annual exceedance probability discharge) and in-channel (100% annual  
 271 exceedance probability discharge) flow conditions across CONUS. This process is diagrammed  
 272 in Figure 2.



273

274 **Figure 2.** Schematic representation of model training for in-channel and bankfull flow top-width  
 275 and depth.

276 2.2.1. Feature Space Reduction

277 Given the large predictor set selected for this study (116), we used two approaches to  
 278 reduce the number of variables while retaining impactful information. This helps increase model  
 279 generalization and interpretability while reducing computational requirements, noise, and

280 dimensionality (Köppen, 2000). The first involves using SHapely Additive Explanations (SHAP;  
281 Lundberg & Lee, 2017) to quantify individual predictor contributions to the target variable.  
282 SHAP values reveal the contribution of each feature to the disparity between the ML model's  
283 prediction and the baseline prediction (ML model's average prediction). Positive values signify a  
284 feature contributes to increasing the model's prediction, while negative values denote the  
285 opposite. The magnitude of the SHAP value indicates the strength of the feature's influence. This  
286 process ranks feature importance, and progressively prunes the least significant predictors in a  
287 series of model retraining and validation cycles, and results in the identification of the balance  
288 between complexity and predictive skill.

289         To address predictor collinearity, we perform correlation analysis following Chan et al.  
290 (2022). We grouped variables based on their correlation and applied Principal Component  
291 Analysis (PCA) to each cluster to create new composite features (Sharma et al., 2015), which we  
292 use as inputs into the model (Figure S2). We specified a target explained variance of 95% to  
293 ensure the retained components collectively account for most of the dataset's variability while  
294 reducing dimensionality (Cruz-Cárdenas et al., 2014). PCA results across the different categories  
295 and their impact on generated components are shown in supplementary materials (Figures S3-  
296 11). We employed the "elbow" heuristic strategy to identify an optimal feature count, which  
297 involves iteratively excluding less informative features while monitoring the coefficient of  
298 determination (R-squared, denoted as ' $R^2$ ') (Liu & Deng, 2020). Upon reaching a plateau in the  
299  $R^2$  value (suggesting the optimal number of features), we identified a subset of 15 features while  
300 the  $R^2$  value remained close to 0.80, signifying these as the most relevant predictors.



### 301 2.2.2. Modeling Channel Geometry

302 The training process involves an out-of-the-box evaluation by training 40 different ML  
303 models, which include neural network, tree-based, and support vector machine approaches to  
304 select top performing models in prediction of top-width and depth as separate models (train on  
305 training data and test on validation data; Figure 2). Then,  $R^2$  and Root Mean Square Error  
306 (RMSE) metrics were used in the objective function to compare the predicted to observed top-  
307 width or depth values, and the 10 top-performing models were selected prior to hyperparameter  
308 tuning (see Table S2). The selected models were then fine-tuned using a 5-fold cross-validation  
309 with 3 repeats and a systematic exploration of hyperparameter combinations. Ultimately, four  
310 different ML models were developed for the estimation of top-width and depth under bankfull  
311 and in-channel conditions.

312 Next, we built a meta-learner and a voting model on the fine-tuned models to harness  
313 collective intelligence. Both the meta-learner and voting model, in this case, use the predictions  
314 of the top 10 fine-tuned selected models, with the exception that a meta-learner has a meta-model  
315 stacked on top to learn from the 10 individual base-level models. These models leverage the  
316 strengths of the top 10 diverse models while minimizing the impact of weaknesses or overfitting  
317 that may occur in any single model. Therefore, we hypothesized that they could be more accurate  
318 than the individual models.

319 Applying this final step helps capture unique insights from the data, reduces overfitting risk,  
320 and enhances overall model performance by leveraging the strengths of the top ten models. The  
321 entire process is conducted on training (70%) and validation (10%) splits of our datasets,  
322 reserving testing (20%) split only for model performance comparison. During the training and  
323 validation phases, the ML configuration with no data transformation or scaling for both

324 predictors and target variables yielded the highest  $R^2$ . This is consistent with the inherent  
325 properties of tree-based algorithms that are non-parametric and base their decisions on splitting  
326 criteria. However, in the case of neural networks, we identified power transformation and scaling  
327 (among the tested log, power, and quantile transformations) as the most suitable techniques  
328 during training.

329 Our out-of-box evaluation showed that XGBRegressor, RandomForestRegressor,  
330 HistGradientBoostingRegressor, LGBMRegressor, ElasticNet, MLPRegressor, BayesianRidge,  
331 ARDRegression, KNeighborsRegressor, and BaggingRegressor were often among the top 10  
332 candidates for our ensemble approach. The highest  $R^2$  was obtained through the meta-learner  
333 ensemble method. A comprehensive list of the models tested, and an example of the outputs used  
334 in the selection process are provided in Table S2. Following the completion of model training,  
335 we utilize the reference fabric aggregated features outlined in section 2.1.4 (Channel Predictors)  
336 to predict channel width and depth under both bankfull and in-channel conditions for  
337 approximately 2.7 million river reaches across the CONUS.

### 338 2.3. Hydrological Traits Impact on Channel Geometry

339 We delve deeper into the impact of features on model performance by examining the top  
340 predictors identified through feature importance analysis. This involves categorizing the values  
341 of significant features into quartiles and assessing the model's skill ( $R^2$ ) in predicting channel  
342 geometry, including channel width and depth under both in-channel and bankfull conditions. By  
343 doing so, we aim to uncover any potential biases of the model towards favoring certain  
344 hydrological characteristics.

345 To assess the impact of varying magnitudes of predictor variables on channel geometry, we  
346 categorized sites based on influential predictors into four quantile ranges: 0-25%, 25-50%, 50-

347 75%, and 75-100% for each predictor. Subsequently, we evaluated how the model's skill ( $R^2$ )  
348 evolves across these quantile ranges, thereby revealing any potential modeling bias towards  
349 specific predictor variables. The presence of discernible patterns in model performance (such as  
350 positive or negative trends from lower to higher quartiles) indicates potential bias in certain river  
351 segments towards either high or low values of the respective attribute. We restricted our analysis  
352 to the four most influential variables identified by SHAP.

353 We utilize SHAP values to comprehend how crucial features influence model outputs (i.e.,  
354 channel width and depth). This involves plotting the SHAP value of a feature against its value  
355 for all examples in the dataset. By doing so, we illustrate how predicted channel dimensions  
356 change as the feature of interest varies, offering insight into complex interactions. This approach  
357 may uncover feature (hydrological) thresholds that influence the model's estimation of channel  
358 depth and width, deviating from simple monotonic relationships between feature values and  
359 channel dimensions. Instead of consistently increasing or decreasing with changes in feature  
360 values, these thresholds may lead to instances where certain increases in feature values result in a  
361 reversal of the model's predicted behavior.

#### 362 2.4. Literature Comparisons

363 Here, we compare the performance of our proposed ML models to other modeling efforts  
364 documented in the literature. The goodness-of-fit metrics employed to compare the proposed  
365 machine learning model with existing literature include the  $R^2$  coefficient (permitting direct  
366 comparison to previous studies), Kling–Gupta efficiency (KGE), and the normalized root mean  
367 square error (NRMSE). We compare our results to that of Blackburn-Lynch et al. (2017), which  
368 introduced regional relations between discharge and channel top-width and depth based on

369 hydrologic landscape regions (HLR; Figure S12; Wolock et al., 2004) and physiographic  
370 provinces. We took the reported  $R^2$  values from Blackburn-Lynch et al. (2017) and juxtaposed  
371 them with those predicted by our ML approach across all test sites, grouped in individual HLR  
372 and Physiographic provinces.

373 Further we compare our model skill to that of global equations founded on drainage area  
374 and discharge, as introduced by Frasson et al. (2019) and Andreadis et al. (2013), respectively.  
375 To facilitate this comparison, we applied these equations to our HYDRoSWOT dataset, using the  
376 total upstream catchment area along with 100 and 50% annual exceedance probability discharge  
377 as input. We also compare our predictions to the recently developed ML method of Doyle et al.  
378 (2023), where we extracted the predicted top-width from their study using the values that are  
379 published in StreamCat data (Hill et al., 2016). Next, we aligned the stream segment IDs  
380 (COMIDs) from their dataset with ours to ensure consistency in comparing stream segments.  
381 Subsequently, we evaluated our predictions of channel top-widths and depths alongside those of  
382 Doyle et al. (2023) against the observed HYDRoSWOT data.

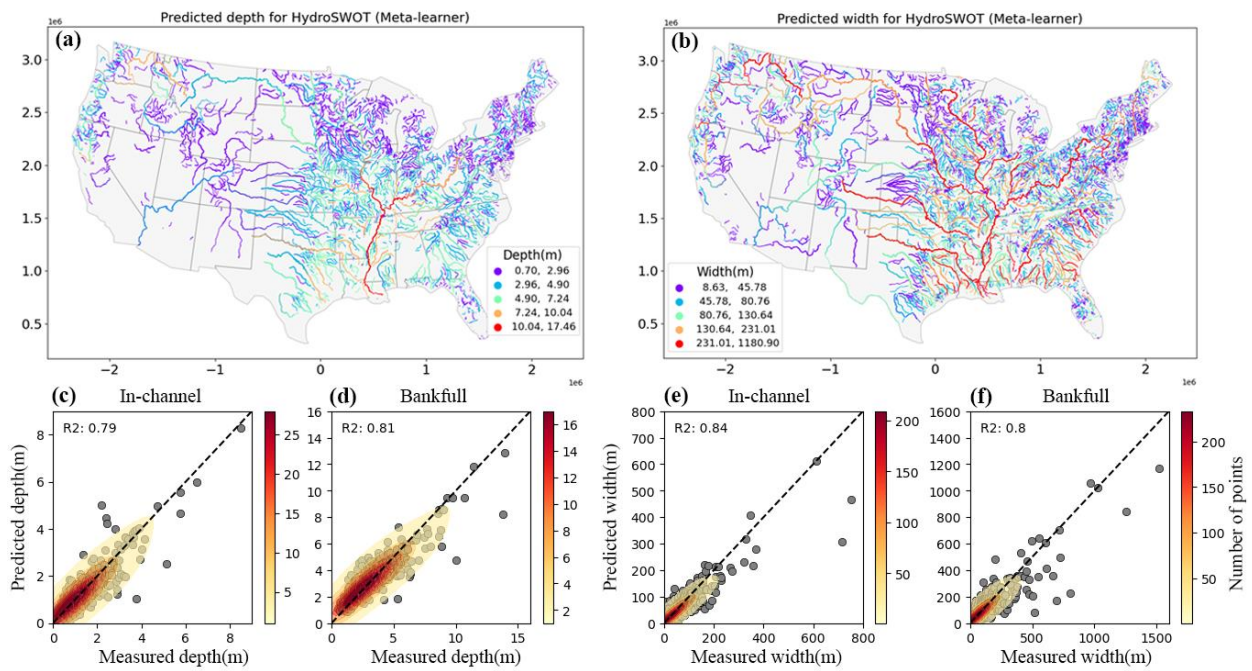
### 383 **3. Results**

#### 384 3.1. In-channel and Bankfull Model Performance

385 For channel top-width, the model  $R^2$  was approximately 0.82 for bankfull and in-channel  
386 conditions, and for channel depth, the model  $R^2$  was approximately 0.80 (Figure 3).

387 Figure 3 shows the performance of all four models on unseen test data. The top-width  
388 models are slightly more accurate (in terms of  $R^2$ ) than the depth models, a trend that is  
389 consistent with previous studies (Booker & Dunbar, 2008). In general, the models underpredict

390 both top-width and depth as channel top-width and depth increase (scatter dots located on the  
 391 right side of the 1:1 line in Figure 3. c-f). This observation can be attributed to the skewed nature  
 392 of hydrologic observations, with proportionately fewer data points collected from larger channels  
 393 (in HYDRoSOT), resulting in less training data for the machine learning model (Krawczyk,  
 394 2016).

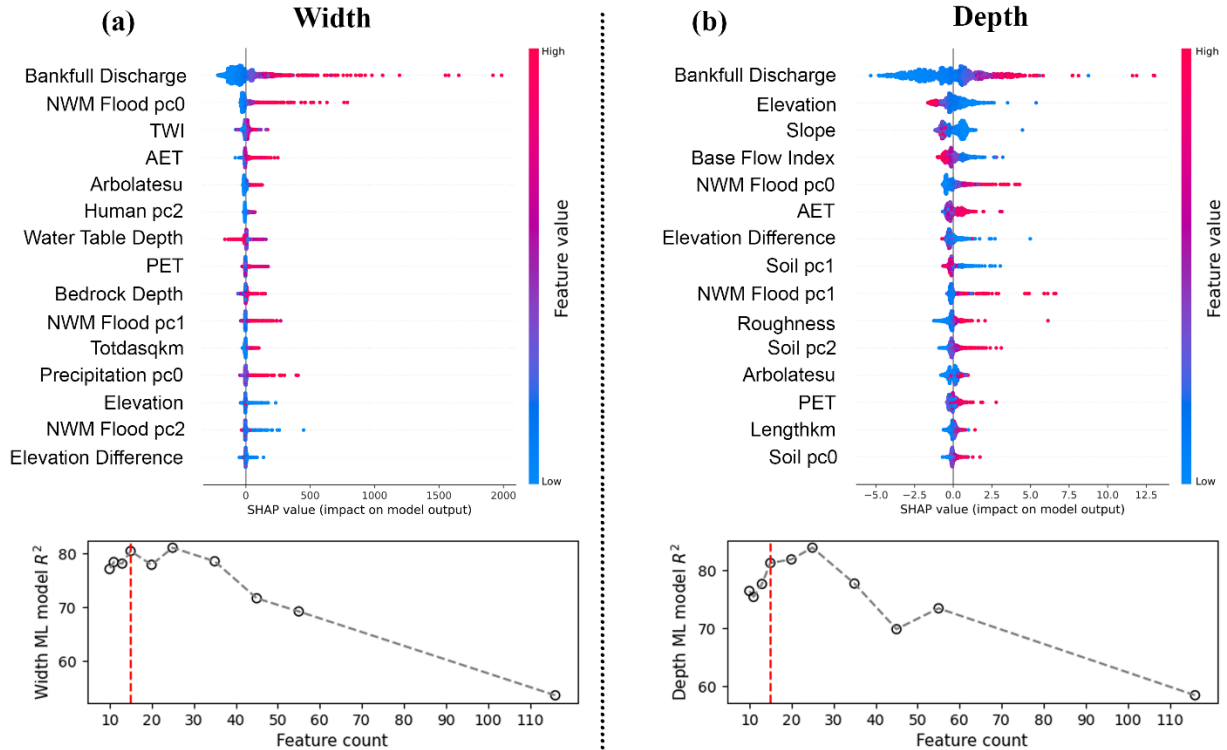


395 **Figure 3.** Predicted bankfull depth and top-width for all used stations in HYDRoSOT data  
 396 mapped to reference fabric flowlines using stream segment IDs (COMID). using meta-learner  
 397 (subplots a and b, respectively). Subplots c and d represent the meta-learner channel depth GOF  
 398 for in-channel and bankfull flow conditions. Subplots e and f represent the meta-learner channel  
 399 top-width GOF for in-channel and bankfull flow conditions.  
 400

### 401 3.2. Impactful Predictors

402 Figure 4 shows the significant features for the prediction of bankfull channel top-width  
403 and depth. It visualizes the SHAP values of each feature across all samples (dots on the plot). It  
404 also arranges features based on the cumulative sum of SHAP value magnitudes across all  
405 samples, and utilizes SHAP values to illustrate the distribution of each feature's impact on the  
406 model output. The most influential feature identified is bankfull discharge and the 0<sup>th</sup> flood  
407 frequency principal component (Figure 4; NWM Flood pc 0). Focusing on top-width (Figure 4a),  
408 the most pivotal variables are the cumulative lengths of all upstream flowlines (Arbolatesu), the  
409 topological wetness index (TWI), and actual evapotranspiration (AET). All three variables  
410 exhibit a positive relationship with channel top-width, as indicated by SHAP importance values.

411 Concerning the feature importance of the depth prediction (Figure 4b), the most  
412 influential variables include bankfull flow estimates, mean catchment elevation, channel slope  
413 (Slope), and the base flow index. Upon comparing the distinctions between the two, it's evident  
414 that channel depth is influenced by topographical factors (elevation and slope), groundwater  
415 recharge (base flow index), and soil texture. Conversely, channel width is influenced by  
416 cumulative stream lengths (Arbolatesu), catchment area (Totdasqkm), TWI, and precipitation.



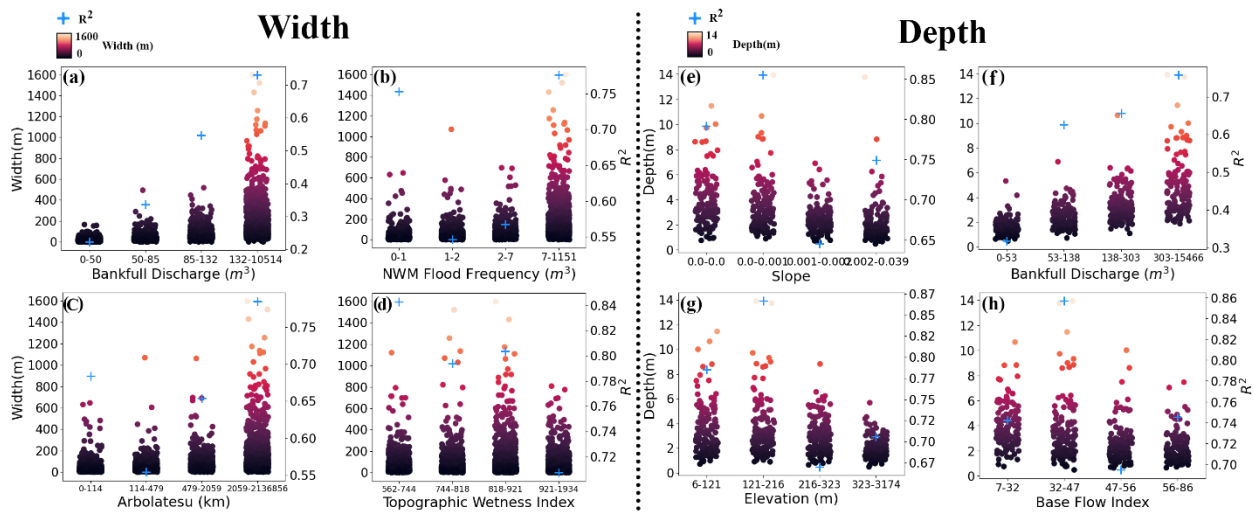
417

418 **Figure 4.** SHAP features importance values, distribution, and impact on model prediction for  
 419 bankfull top-width (a) and depth (b). All NWM Flood principal components (PCs) are  
 420 uncorrelated representations (PCs are orthogonal to each other) of 20-0.1% annual exceedance  
 421 probability discharge. Human PC2 is a representation of the total reservoir volume in the  
 422 catchment. Soil PC0 represents soil moisture, Soil PC1 represents % of clay and silt, and Soil  
 423 PC2 represents % of sand.

424 **3.3. Impact of Predictors Magnitude on Model Performance**

425 Next, we examined the four most influential features in each model by partitioning their  
 426 values into quartiles, allowing us to assess model performance across segments as shown in  
 427 Figure 5. In the case of channel top-width, both bankfull discharge and the flood frequency PC0  
 428 show similar patterns when assessing model performance (Figure 5a, b), showing that as channel  
 429 top-width increases, modeling errors decrease. A pattern was also found in the arbolate sum  
 430 (Arbolatesu; Figure 5c). The topographic wetness index (TWI; Figure 5d) values do not appear  
 431 to have a strong pattern between quartiles and performance.

432 Looking at the depth model in Figure 5e, we see that as the channel slope decreases,  
 433 model GOF ( $R^2$ ) increases. Model performance declines when bankfull discharge is smaller  
 434 (Figure 5f). Higher elevations (Figure 5g) also tend to negatively impact model GOF ( $R^2$ ), and  
 435 the model showcases its highest performance in regions dominated by runoff rather than  
 436 groundwater contributions (baseflow index; Figure 5h).



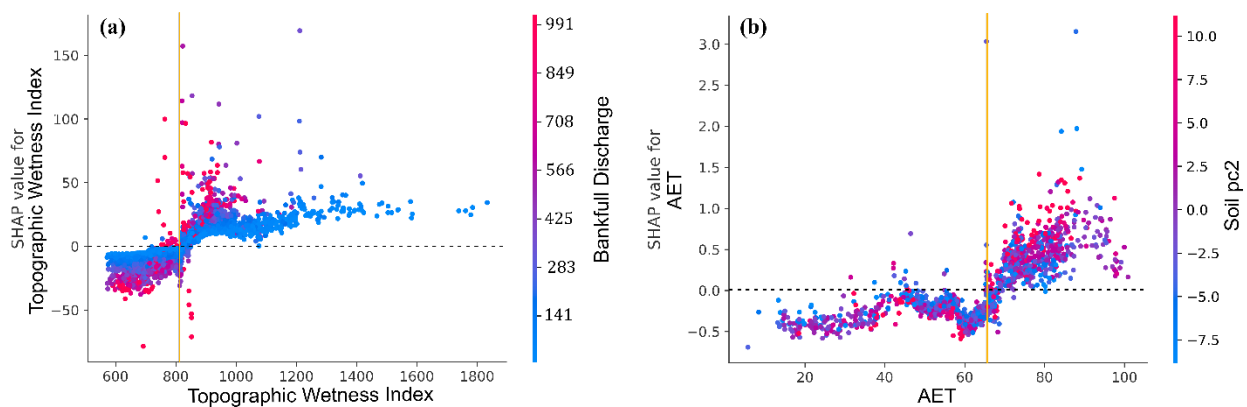
437 **Figure 5.** ML model performance across different quantiles of the most influential variables for  
 438 both depth and width models. Within each subplot each dot represents a station belonging to a  
 439 quantile of an important feature a, b, c, and d belong to grouped performances of top-width and  
 440 subplots e, f, g, and h belong to depth. Th model skill ( $R^2$ ) is shown as cyan plus sign.

442 3.4. Hydrological thresholds

443 It has been theorized that channel geometry follows stable hydrologic regimes (Rosgen,  
 444 1994). To elucidate this, we investigate the interrelation between key variables and their  
 445 respective influence on model prediction, as depicted in Figure 6. In Figure 6a, the significance  
 446 of the Topographic Wetness Index (TWI) in forecasting channel top-width is quantified. The



447 absolute magnitude of the SHAP value serves as an indicator of the importance of each site (dots  
448 on the plot) in machine learning. A positive SHAP value suggests an increase in channel  
449 dimension, while a negative value indicates the opposite. The illustration reveals that channels  
450 with elevated bankfull discharge (depicted by purple dots) exhibit particularly informative data,  
451 characterized by high positive or negative SHAP values, during the machine learning model's  
452 training phase. Interestingly, there is an inflection point near  $TWI = 810$ . Here, high discharges  
453 (pink dots) linked with  $TWI < 810$  result in a decrease in channel top-width (indicated by  
454 negative SHAP values), while values above  $TWI > 810$  correspond to an increase in channel top-  
455 width. This is also observed when looking at bankfull discharge high (purple) and low (blue)  
456 values, as higher discharge values do not linearly contribute to greater channel dimensions. This  
457 suggests that channel morphology shifts after  $TWI = 810$  and is most evident in larger channels  
458 (or higher discharge/pink dots). Similarly, an inflection point of  $AET = 65$ , as shown in Figure  
459 6b, illustrates that areas characterized by  $AET < 65$  contribute to shallower channels, and  $AET$   
460 greater than this threshold is an indicator of deeper channels. The soil PC2, which is a  
461 representative of clay and silt components (see supplementary materials in Figure S3), is evenly  
462 scattered and shows no pattern.

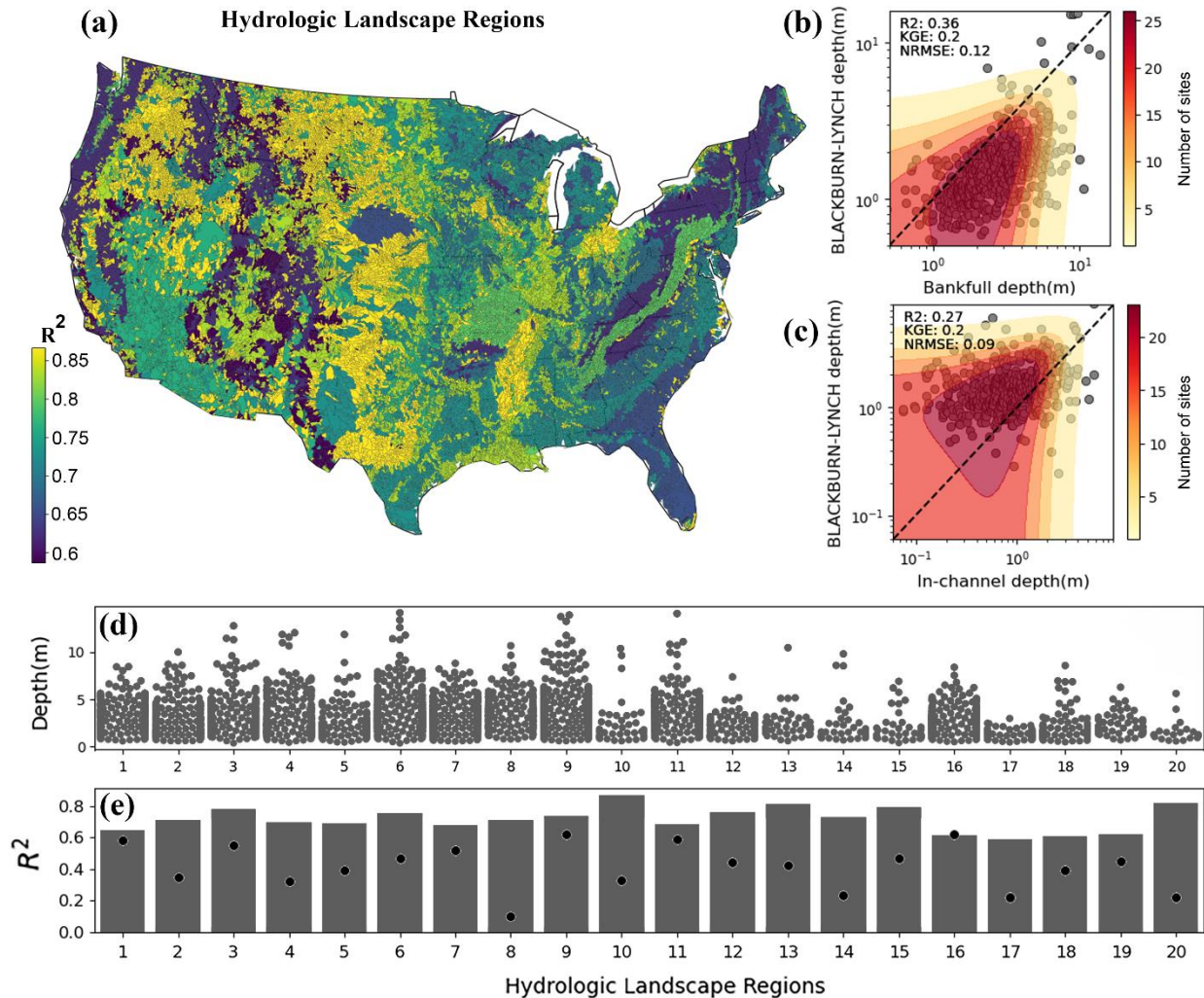


464 **Figure 6.** Representation of the contribution of each important feature and its inner correlation  
465 and impact on the top-width and depth models. (a) represents the impact of the topological  
466 wetness index on the channel top-width ML model and its correlation with the NWM 50%  
467 annual exceedance probability discharge. (b) represents the impact of AET on the channel depth  
468 ML model and its correlation with % of soil clay and silt content.

### 469 3.5. Performance in Hydrologic Landscape Regions

470 . It reveals a pattern where channel depth models show lower GOF ( $R^2$ ) in the prediction  
471 of bankfull depth in regions with impermeable bedrock (HLR 16, 17, 18, and 19; Carlier et al.,  
472 2018; Santhi et al., 2008). Figures 7b and c show that the GOF of regional drainage area-based  
473 equations provided by Blackburn-Lynch et al. (2017) over the CONUS is poor ( $R^2 \sim 0.2-0.3$ ).  
474 We took the reported  $R^2$  values from Blackburn-Lynch et al. (2017) and compared it to the one  
475 predicted by the proposed ML approach in Figures 7d and e. Figures 7d and e, highlight the  
476 potential correlation between model performance, the distribution of the target variable, and the  
477 quantity of available points for machine learning. Results showed that the proposed depth model  
478 shows a significant improvement in  $R^2$  across all HLR and physiographic provinces (see  
479 supplementary Figure S13; except for physiographic provinces 1 (Adirondack) and 23 (Southern  
480 Rocky Mountains)). In hydrological regions, HLR 16, characterized as humid mountains with  
481 permeable soils and impermeable bedrock, the ML model did not show improved GOF  
482 compared to the Blackburn-Lynch et al. (2017) study ( $R^2 \cong 0.60$  of our model vs.  $R^2 \cong 0.62$   
483 reported). Moreover, regions with impermeable bedrock have lower range of measured channel  
484 depths (HYDRoSWOT measurements; as depicted in Figure 7d). Furthermore, we provide the

485 improvements of ML-based approaches over discharge-based regional equations by Bieger et al.  
486 (2015) in supplementary information (Figure S13).

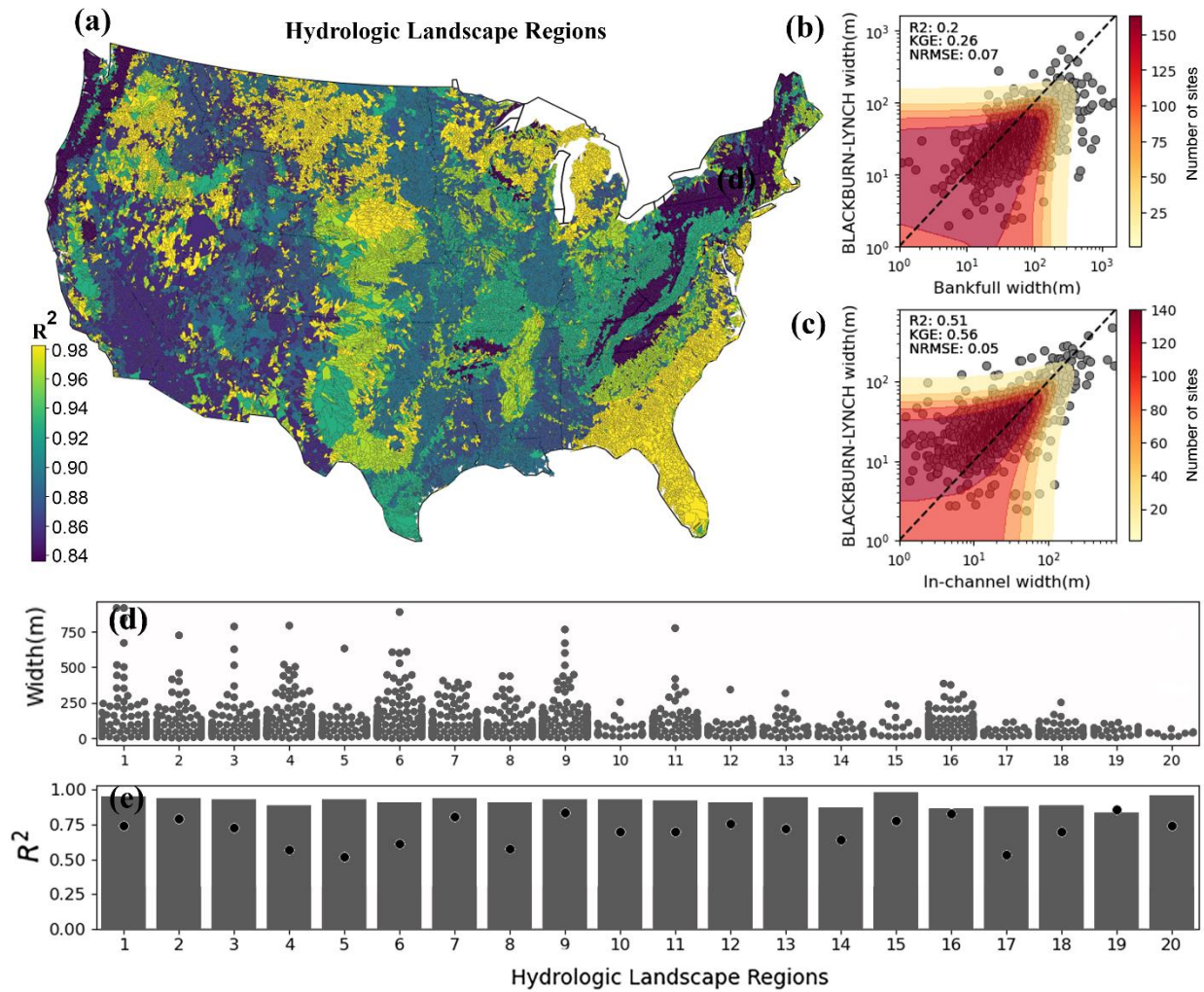


487

488 **Figure 7.** Depth ML model GOF ( $R^2$ ) comparison with literature and across hydrologic  
489 landscape regions (a). (b) the GOF ( $R^2$ ) of Blackburn-Lynch et al. (2017) derived depth from  
490 NWIS bankfull depth measurements. (c) GOF ( $R^2$ ) of Blackburn-Lynch et al. (2017) derived  
491 depth from NWIS in-channel depth measurements. (d) shows the population of sites and  
492 distribution of depth per HLR in HYDRoSOT. (e) proposed ML model GOF ( $R^2$ ) comparison

493 (in columns) to Blackburn-Lynch et al. (2017) (in black dots) across hydrologic landscape  
494 regions.

495 Performing the same operations for the top-width ML model (Figure 8a) and comparing model  
496 skill ( $R^2$ ) to Blackburn-Lynch et al. (2017) (Figure 8b and c), we see an overall higher  $R^2$   
497 performance in the tested ML model (Figure 8d and e and Figure S14). Looking at HLRs, we see  
498 that the lowest GOF comes from region 19, which is described as very humid mountains with  
499 permeable soils and impermeable bedrock with  $R^2 \cong 0.75$ . The highest GOF based on  $R^2$  are seen  
500 in all plains (HLR 1-8) and plateau regions (i.e., HLR 9-13; Vidon & Hill, 2004). All these  
501 plateaus are categorized as surface flow response (in contrast to groundwater flow response)  
502 regions (Wolock et al., 2004). On the other hand, the mountainous regions in HLR 16-19 have  
503 the lowest GOF (in terms of  $R^2$ ). Compared to depth estimates, the drainage area-based  
504 equations by Blackburn-Lynch et al. (2017) have a better skill in the prediction of top-width  
505 under in-channel conditions (Figure 8c;  $R^2 \sim 0.5$ ) but are not appropriate for bankfull flows  
506 (Figure 8c;  $R^2 \sim 0.2$ ).

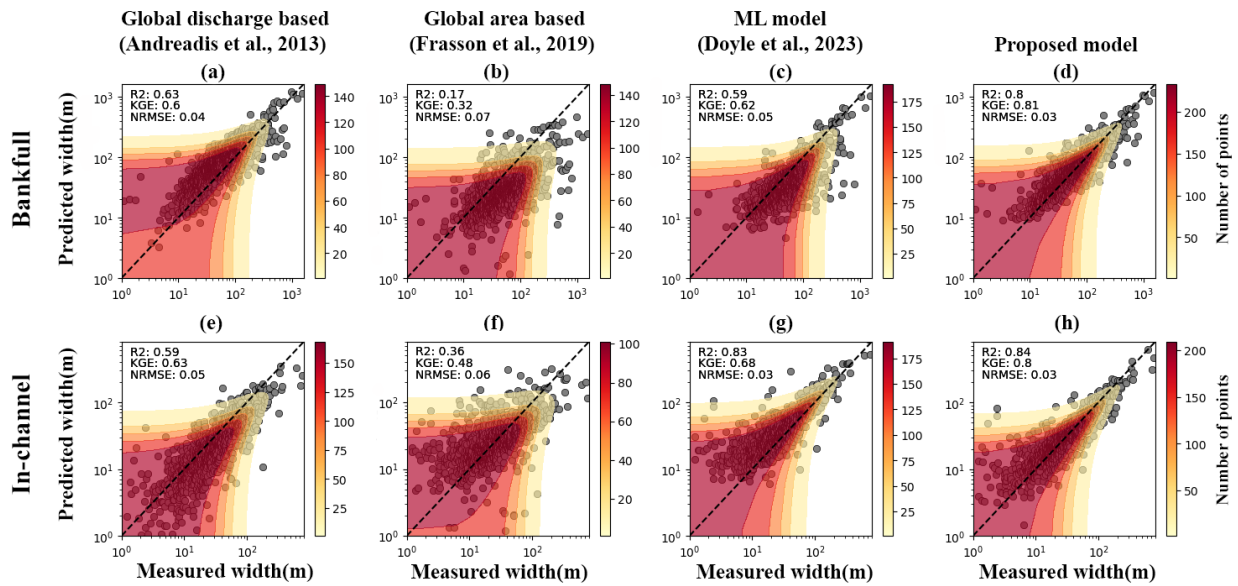


507

508 **Figure 8.** Width ML model GOF ( $R^2$ ) comparison with literature and across hydrologic  
 509 landscape regions (a). (b) the GOF ( $R^2$ ) of Blackburn-Lynch et al. (2017) derived width from  
 510 NWIS bankfull width measurements. (c) GOF ( $R^2$ ) of Blackburn-Lynch et al. (2017) derived  
 511 width from NWIS in-channel width measurements. (d) shows the population of sites and  
 512 distribution of width per HLR in HYDRoSWOT (e) proposed ML model GOF ( $R^2$ ) comparison  
 513 (in columns) to Blackburn-Lynch et al. (2017) (in black dots) across hydrologic landscape  
 514 regions.

515 3.6. Comparative Model Assessment

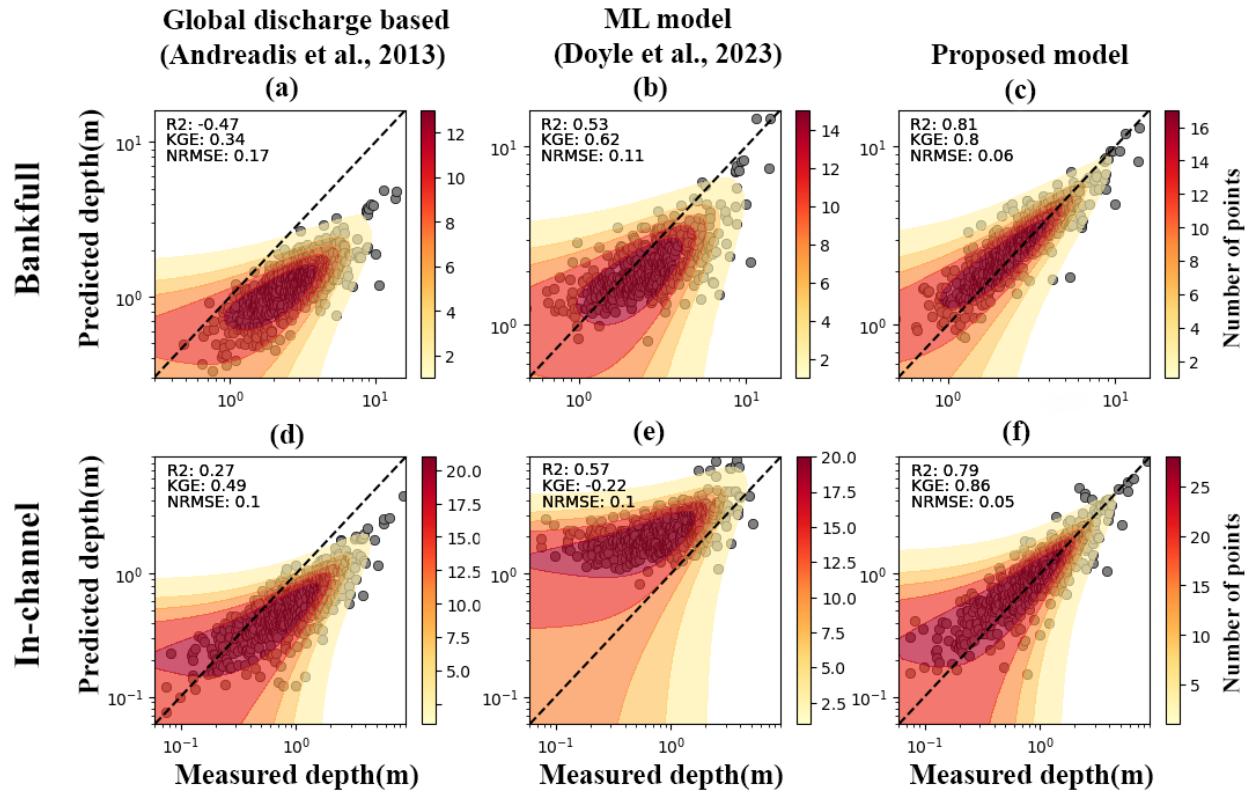
516 Following the substantial enhancement in performance when compared to regionally  
517 based discharge equations, we compare our results to global discharge and drainage area based  
518 questions and recent ML model by Doyle et al. (2023). Figure 9. shows the comparison of the  
519 global models and the recently proposed ML-based approaches, where ML models show a clear  
520 superiority over global relations.



521  
522 **Figure 9.** Channel top-width at bankfull and in-channel flow conditions model performance  
523 comparison between (a, e) global drainage area-based equations, (b, f) global discharge-based  
524 equations, (c, g) the ML model proposed by Doyle et al. (2023), and (d, h) our proposed model.

525 Using the same process, we can derive and compare our proposed depth model to other  
526 models in the literature (Figure 10).





527

528 **Figure 10.** Channel depth at bankfull and in-channel flow conditions model performance  
 529 comparison between (a, e) global drainage area-based equations, (b, f) global discharge-based  
 530 equations, (c, g) the ML model proposed by Doyle et al. (2023), and (d, h) our proposed model.

531 With improved representation of discharge characteristics (from NWM), the proposed  
 532 ML model has higher GOF (in terms of  $R^2$ ) compared to its predecessors.

### 533 4. Discussion

#### 534 4.1. Significance of Discharge in Channel Geometry Modeling

535 The outcomes of the SHAP features importance analysis suggest that the machine  
 536 learning models have effectively captured the variations in channel dimensions while preserving

537 the faithful representation of diverse hydrological and hydraulic processes and their  
538 interrelations. Conventional methods often employ empirical equations that relate these  
539 dimensions to drainage areas (Dunne & Leopold, 1978). More recent studies have focused on the  
540 application of bankfull discharge rather than drainage areas for developing channel dimension  
541 relations (Bieger et al., 2015). In Figure 6a we saw a nonlinear relationship between bankfull  
542 discharge/TWI and channel dimensionality. This non-linear relation was also reported by  
543 Erikson et al. (2024) recent study that showed regions exhibiting non-linear discharge scaling,  
544 bankfull channel dimensions increase more rapidly with drainage area compared to areas with  
545 linear discharge scaling. This suggests that the recurrence interval of the characteristic discharge  
546 determining channel geometry may be greater in regions with non-linear discharge scaling than  
547 in those with linear scaling.

548         We found the most influential feature identified for both depth and width is bankfull/in-  
549 channel discharge and is aligns well with the feature importance analysis conducted by Doyle et  
550 al. (2023) for wetted top-width and thalweg depth, as well as studies where authors employed in-  
551 situ measurements to establish relationships between discharge and channel characteristics using  
552 satellite imagery (Bjerklie, 2007; Bjerklie et al., 2005; Zakharova et al., 2020). Our findings  
553 underscore the significant role of using 50% annual exceedance probability discharge derived  
554 from NWM retrospective data when NWIS-derived 50% annual exceedance probability  
555 discharge is unavailable (the most important predictor shown Figure 4). The usage significantly  
556 enhances the GOF of prediction of channel top-width and depth, as indicated by the increase in  
557  $R^2$  value. It consistently (in width and depth under both flow conditions) demonstrates a strong  
558 positive correlation between higher magnitudes of the flood (Bankfull discharge and NWM



559 Flood PCs), and broader, deeper channels, aligning with the established literature (Bjerklie et al.,  
560 2005; Brown & Pasternack, 2014; Wohl & Wilcox, 2005).

561           Interestingly, global equations based on drainage area show significantly lower GOF  
562 ( $R^2$ ), while methods grounded in discharge demonstrate notably higher levels of GOF. However,  
563 the use of hydraulic geometry curves does require a comprehensive understanding of the  
564 bankfull or in-channel discharge at specific locations of interest. To overcome this issue, Doyle  
565 et al. (2023) used the mean runoff from RunoffWs (StreamCat data; McCabe & Wolock, 2011)  
566 that is generated based on a 4 km × 4 km water-balance model for a period of 1900–2008. We  
567 improved upon this by replacing the water balance model with more accurate NWM  
568 retrospective data and derived various discharge characteristics by computing discharges of 20-  
569 0.1% annual exceedance probability and long-term average flow percentiles from NWM v2.1  
570 historical archives (Michael et al., 2023). Lin et al. (2020) also identified discharge as the  
571 primary significant characteristic both in bankfull top-width and depth estimations. However,  
572 Doyle et al. (2023) identified the watershed area as the primary significant feature for their  
573 models in feature importance analysis. Contrarily, in our study, we demonstrated that a more  
574 accurate discharge from NWM (compared to Doyle et al.'s (2023) water balance model) can  
575 result in higher model GOF and the identification of discharge as the primary significant feature.

#### 576           4.2. Other influential variables in Channel Geometry Modeling

577           Other most influential variables for the modeled depth were identified to be channel slope  
578 (Slope), mean catchment elevation (also reported by Lin et al. (2020) as the third important  
579 variable), and the base flow index. All these variables demonstrate a negative relationship with  
580 channel depth, suggesting that regions characterized by lower elevations and mild slopes (e.g.,

581 coastal areas) combined with lower groundwater to runoff contribution (low BFICat values)  
582 favor deeper channels (Biswal et al., 2023; Harman et al., 2008). These insights contribute to a  
583 comprehensive understanding of the multifaceted relationships between the examined features  
584 and channel dimensions (Morel et al., 2020).

585 In the case of width models, other influential variables were namely Arbolatesu, TWI,  
586 and AET that showed a positive relationship with channel top-width, as indicated by SHAP  
587 importance values. Practically, this can be interpreted to mean that regions with a high  
588 Topological Wetness Index, higher vegetation density and coverage (AET), or extended  
589 upstream drainage area tend to possess broader channels.

#### 590 4.3. Utility and Limitations

591 The utility of bankfull top-width and depth estimates is evidenced in a broad body of  
592 literature and their application in large-scale modeling (such as the National Water Model) and  
593 observation (such as the Surface Water and Ocean Topography (SWOT) mission). Across the  
594 hydrological and hydraulic sciences, these estimates serve applications across studies  
595 encompassing research, engineering, modeling, and mapping applications. As model domains  
596 and resolutions expand, the capability to precisely estimate these attributes within high-  
597 resolution and dynamic hydrographic networks is becoming more important and useful  
598 (Archfield et al., 2015; Wood et al., 2011). Despite advancements in remote sensing applications  
599 for monitoring, the current resolution do not yet align with the necessary level of detail. Hence,  
600 there persists a demand for modeling and estimating these characteristics across diverse and  
601 heterogeneous landscapes.

602 To date, there have been efforts to generalize these relationships with respect to  
603 hydrographic (drainage area) and landscape (StreamCat) traits. These efforts encompass global  
604 and regional regression models, along with the application of random forest techniques across an  
605 extensive array of predictor variables at the catchment level. In this work, we aimed to enhance  
606 prior approaches by introducing two distinctive augmentations.

607 The initial addition was to incorporate comprehensive hydrographic details and historical  
608 streamflow model simulations from NWM aggregated based on reference fabric apart from other  
609 landscape attributes and drainage area. This establishes a dynamic, updatable channel geometry  
610 dataset for United States reference flow (Blodgett et al., 2023), notably applied in numerous  
611 integrated federal modeling initiatives. Consequently, our findings underscored elevation, slope,  
612 arbolate sum (Arbolatesu), drainage area, and length as pivotal predictors in estimating in-  
613 channel and bankfull flow conditions. The aggregation scale of these predictors emphasizes the  
614 potentially apparent yet significant role of network configuration in shaping bankfull flow  
615 conditions. Moreover, the incorporation of streamflow statistics derived from a comprehensive  
616 hydrologic model (National Water Model 2.1) demonstrated notable predictive power in  
617 establishing bankfull conditions. While Figure 1 highlights the level of agreement between the  
618 one-year ( $R^2 = 0.3$ ) and two-year ( $R^2 = 0.79$ ) flow conditions with the observed record, it is  
619 evident that the patterns in the data were able to help the ML model deduce more accurate  
620 estimates (based on  $R^2$ ) across the entire network.

621 The subsequent addition to this research was to introduce a more robust ML method to  
622 address this challenge. We minimized the complexity of the model by retaining only the most  
623 impactful predictors. This process helped reduce the number of model predictors from 116 to 15.  
624 This represents a major improvement over previous efforts such as Doyle et al. (2023), who

625 performed a similar test but concluded with 96 predictors, many of which exhibit high  
626 collinearity. Also, diligent attempts were undertaken to minimize data covariance and refine the  
627 input parameter space. A comprehensive evaluation encompassed a broad spectrum of "out-of-  
628 the-box" (40) and fine-tuned models (10), culminating in the development of a meta-learner  
629 model that harnessed the expertise and diversity of the collective models. These efforts yielded  
630 four models that exhibited enhanced predictive capabilities compared to existing methodologies.  
631 Beyond model development, this approach was applied to the entire NHGF network, resulting in  
632 outcomes integrated into the core data product. This integration actively improves data, making  
633 it more findable, accessible, interoperable, and reusable within the expanding hydrologic science  
634 data system.

635         Despite the numerous advantages, certain limitations warrant consideration. One  
636 limitation is the training data and its general applicability, which are constrained to USGS sites  
637 only. These sites are often situated in locations with high banks that confine the flow for ease of  
638 measurement and, as such, may not precisely represent the geometry of the entire reach. Also in  
639 estuaries, particularly near the river mouth where it widens into the coastal zone, there are no  
640 recorded ADCP measurements in HYDRoS<sub>W</sub>OT, and model accuracy is unreliable. This  
641 limitation in training dataset also result in unreliable estimates near manmade structures (e.g.,  
642 dams and ponds). Another limitation stems from the primary significant feature of ML models,  
643 or the 100% annual exceedance probability discharge derived from the NWM, where the NWM  
644 diminished skill (as shown in Figure 1) in this exceedance probability, impacting the overall  
645 model GOF. While this method offers consistent reach-averaged channel geometry data for the  
646 entire CONUS, its resolution is constrained to the reach scale. The consequences of averaged  
647 reach-scale resolution can vary depending on the application. On one hand, it may lead to

648 limitations such as overlooking intricate channel features like braiding, incision, and aggradation.  
649 On the other hand, it can offer benefits such as simplifying complexity and reducing  
650 computational efforts and can be fused with DEM derived products to represent the entire river  
651 reach. Achieving a resolution lower than reach scale would necessitate the utilization of satellite  
652 imagery for accurate width estimation. However, even with satellite imagery, resolving depth  
653 accurately at resolutions lower than reach scale remains a challenge in the absence of  
654 bathymetric related information.

655           Moreover, although our modeling strategy effectively reduced 112 variables into 15  
656 components, the reliance on source data restricts the scalability of this approach to different  
657 geographical locations. Substantial data challenges associated with using existing NWM  
658 retrospective data (Johnson J. Michael et al., 2023) and its sheer size currently make this dataset  
659 globally unavailable, although less accurate global hydrological models can be substituted.  
660 Consequently, our model's applicability could be constrained in global regions where a  
661 comprehensive long-term historic modeled record is unavailable. The ML modeling method's  
662 applicability extends beyond a singular regression relation, thereby potentially posing challenges  
663 in its adaptation to new networks due to the complexity of the parameter set and model  
664 deployment. Nonetheless, envisioning future applications within the United States, the  
665 comprehensive coverage across our domain, coupled with the seamless integration into the  
666 continually evolving data system, holds significant promise for advancing the evolution and  
667 dissemination of hydrofabric data.

668           Anticipating advancements within the United States Next Generation Water Resource  
669 Model initiative led by NOAA and Cooperative Institute for Research to Operations in  
670 Hydrology (CIROH), the availability of this information will certainly help many ongoing

671 efforts. Notably, the team spearheading the development of routing infrastructure for the  
672 NextGen system relies on approximations of general and natural cross-sectional data to facilitate  
673 Muskingum-Cunge and Diffusive Wave routing methods. The first of these can be improved or  
674 derived from a quality estimate of bankfull top-width and depth using existing methods like  
675 those in WRF-hydro or other channel geometry estimators (e.g., parabolic forms such as in  
676 Dingman, 2009). Additionally, the Flood Inundation Mapping team can incorporate estimated  
677 depths into their synthetic rating curve workflow (Johnson et al., 2019; Zheng et al., 2018),  
678 addressing the current limitations in accounting for bathymetry within DEM-derived hydraulic  
679 states. The inclusion of an estimated bankfull depth can help better estimate the volume of water  
680 in the channel and reduce instances of systematic over- or under-prediction (Johnson et al.,  
681 2019). While the effectiveness of these diverse use cases awaits validation, their accessibility  
682 through fundamental geospatial products supporting USGS, NOAA, and CIROH efforts will  
683 facilitate simplified testing and evaluation processes.

## 684 **5. Conclusions**

685         The measurements of channel top-width and depth, as well as their ratio, particularly at  
686 bankfull flow conditions, play a pivotal role in the fields of hydrology and river science  
687 (Bjerklie, 2007). These metrics hold profound importance for several key reasons. First and  
688 foremost, they are fundamental to characterizing the physical attributes and morphology of river  
689 systems, providing critical insights into the geometry of channels. (Luo et al., 2007)

690         Also, NASA's Surface Water and Ocean Topography (SWOT; Biancamaria et al., 2016)  
691 mission objectives, which encompass the precise measurement of channel discharge, water  
692 surface elevation, and variations in channel dimensions, rely on the calibration and validation of

693 its remote sensing instruments, and the proposed method could play a pivotal role in ensuring the  
694 mission's accuracy and its valuable contributions to our comprehension of Earth's surface water  
695 dynamics. In smaller rivers and tributaries, estimating channel dimensions presents a  
696 considerable challenge, primarily due to the restricted spatial resolution of satellite data.  
697 Nonetheless, as showcased in this study, ML approaches trained on hydrographic information  
698 and model simulated discharge values provide an alternative method to adequately capture  
699 channel dimensions.

700 Our study showcases the efficacy of the current ensemble ML model in minimizing the  
701 necessary predictors while mitigating concerns related to multicollinearity and model confusion.  
702 Leveraging the extensive dataset of NOAA's Office of Water Prediction's (OWP) NWM flow  
703 characteristics encompassing 2.7 million reaches, we have achieved notably high GOF (reported  
704 as  $R^2$ ) in predicting bankfull depth and top-width. Nonetheless, it is pertinent to acknowledge  
705 that the model's GOF diminishes with reduced discharge levels. This result again points to the  
706 challenge of getting equal skill in large and small segments of the network. This pronounced  
707 effect in smaller rivers and tributaries suggests a potential association with NWM's limited skill  
708 in representing base flow, a key determinant of channel dimension within ML models (as shown  
709 in Figure 5, there is a correlation between lower discharge and reduced model GOF). Certain  
710 constraints include the exclusive use of training data from USGS sites, primarily situated in areas  
711 with elevated banks, restricting flow and potentially not faithfully representing the complete  
712 reach geometry. We conducted extensive model evaluations across different hydrologic  
713 landscape regions and physiographic provinces and divisions. Our comprehensive analysis  
714 substantiates the capacity of our approach to augment the existing models documented in  
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724 **Data Availability**

725 The datasets used in this study, including the StreamCat data (Hill et al. 2016), Reference Fabric  
726 (Bock et al., 2022), HYDRoSWOT (Canova et al., 2016), NWM (Johnson et al., 2023), and  
727 ClimateR (Johnson & Clarke, 2019) are publicly available.

728 The ML predictions for channel depth and width are available at:

729 Modaresi Rad, A. (2024). Machine Learning-Derived Channel Width and Depth for the National  
730 Hydrologic Geospatial Fabric in CONUS, HydroShare,

731 <http://www.hydroshare.org/resource/d147fcf554a54b2aaa4f146f85da0e03>

732 The developed code for this research are available at:

733 [https://github.com/LynkerIntel/bankfull\\_W\\_D](https://github.com/LynkerIntel/bankfull_W_D)

734

735 **Disclaimer**



736 The views expressed in this article do not necessarily represent the views of NOAA or the United  
737 States.

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